

Vesiculation of a rhyolitic melt

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1. Introduction

- Bubble growth in rhyolitic melts is a primary control on some of the largest explosive eruptions. However, vesiculation remains poorly constrained.
- Few studies have captured in-situ vesiculation of a rhyolitic melt, but have rather relied on interpretation of quenched natural or experimental samples. The previous in-situ study^[1] followed vesiculation in water-poor ($^{\circ}$ 0.14 wt%) rhyolitic melt at P = 1 Atm.
- This work aims to provide measurements of rapid insitu vesiculation in more water-rich rhyolitic melt from high resolution imagery. Results are relevant to post-fragmentation magma vesiculation.

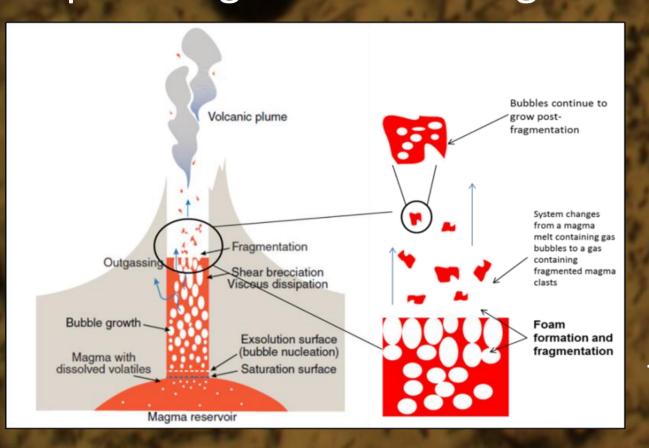
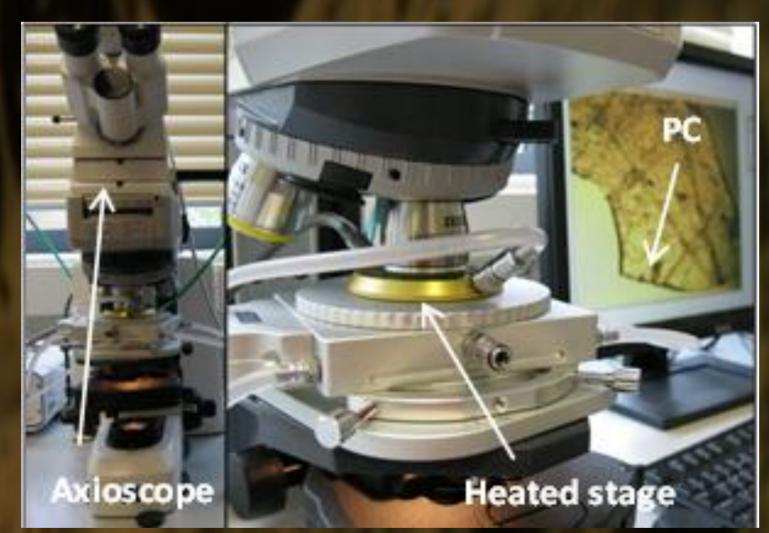


Fig 1. The stages of bubble nucleation and growth leading to magma fragmentation in a Bubbles volcanic conduit. continue atmospheric pressure fragmentation. Modified from

2. Methods

- Using the technique of Applegarth et al (2013), thin wafers ($\sim 100 \, \mu m$ thick) of obsidian (0.97 wt% H₂O) from the 2008 eruption at Chaitén, Chile were held from 5 minutes up to 2 days in the hotstage at between 575 °C and 875 °C. All experiments were conducted at 1 atm and therefore do not consider growth by decompression.
- The in-situ growth of many individual bubbles were recorded directly to PC and then measured using particle tracking code written in MATLAB.

Fig 2. Hot-stage microscope setup³. Sample is placed inside a ceramic furnace (Linkam TS1500 heated stage), mounted on a Zeiss Axioscope.



The potential for sample dehydration was considered by estimating the extent of diffusive degassing from wafer surfaces using simple diffusion models^[4]. Dehydration was found to be negligible during brief high temperature experiments but became increasingly important for slower, lower-temperature experiments

3. Physical processes of bubble growth

- Five stages of bubble growth were directly observed (Fig. 4).
- Most rapid average bubble growth rate at $875 \,^{\circ}$ C (1.27 μ m s⁻¹; $\eta = 10^{7.1}$ Pa s)
- Slowest bubble growth rate at 725 °C (0.02 μ m s⁻¹; $\eta = 10^{9.20}$ Pa s)
- No bubble growth was noticeable below 725 °C.
- Growth rates decreased with time, as reported in [1]. However, growth rate decreases observed here are due to bubble-bubble interactions.

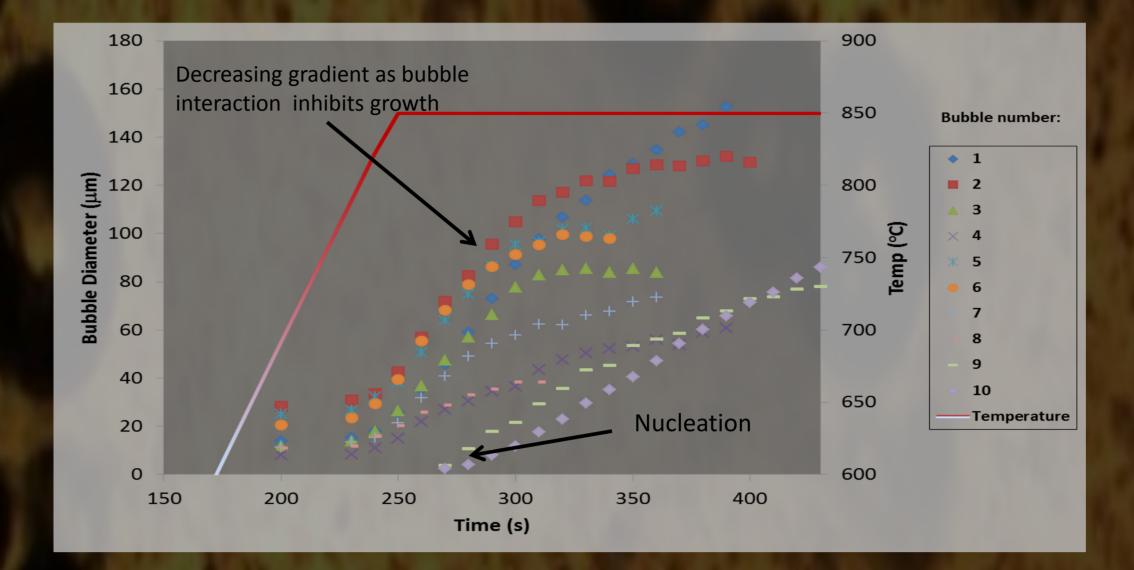
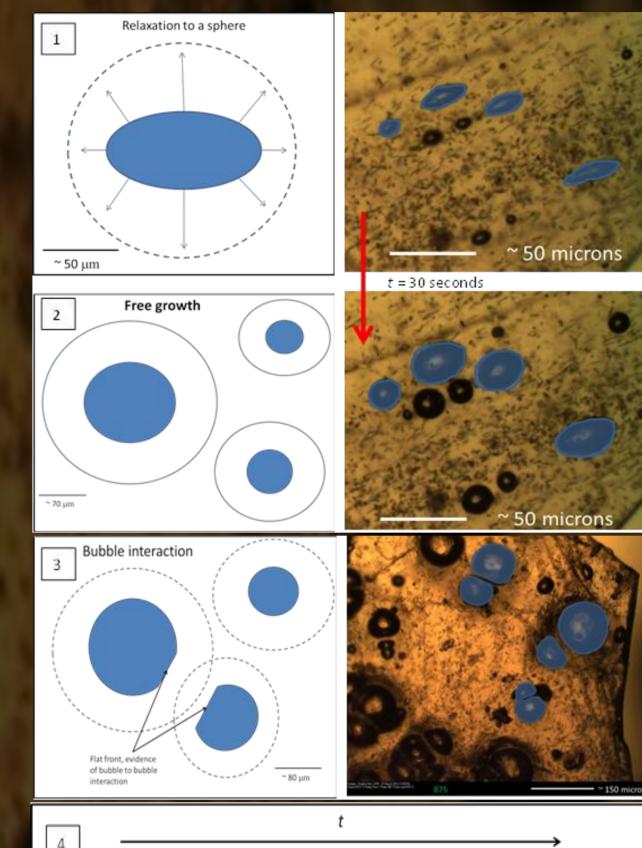


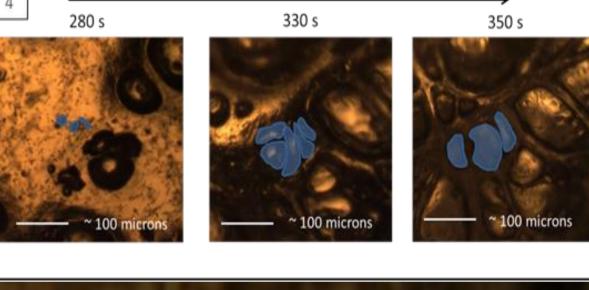
Fig 3. Bubble growth over a period of > 200 seconds at temperatures of 850 °C and 775 °C.

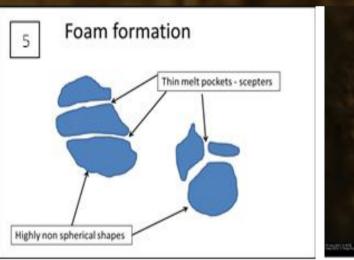
Experimentally obtained bubble growth rates were compared to the predicted growth rates modelled by Navon et al (1998), Fig. 6, using: $P_g - P_f = \frac{2\sigma}{R} + 4\eta \frac{V_r}{R}$

Where the bubble overpressure (Pg) - ambient pressure (P_f) is assumed to be initially constant, and the effect of surface tension (σ) is ignored for bubble radius > 5.5 μ m :. Growth rate $(V_r) = R/4\eta$

- Viscosity (η) was estimated using the model of [7].
- At all temperatures modelled growth rates exceed those determined experimentally; there is a closer fit at higher temperatures.
- This finding is contrary to [1], whose modelled growth rate fitted better at lower temperatures. This discrepancy may relate to the varied water contents of samples studied.







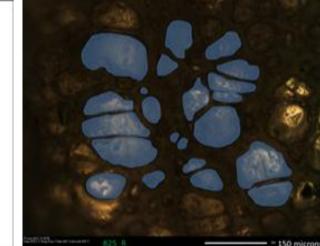


Fig 4. Bubble growth morphology

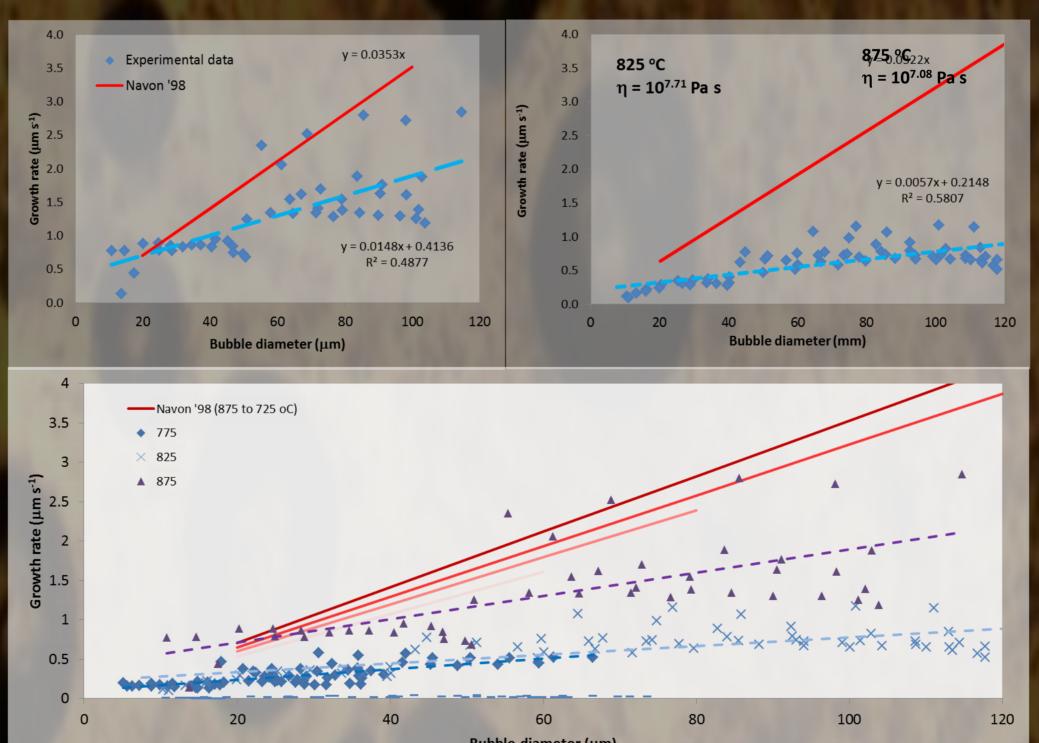


Fig 5. Comparison of experimental bubble growth rates compared to those modelled by Navon et al, 1998.

Using the techniques described it is also possible to track bubble size distributions and bubble number densities through time.

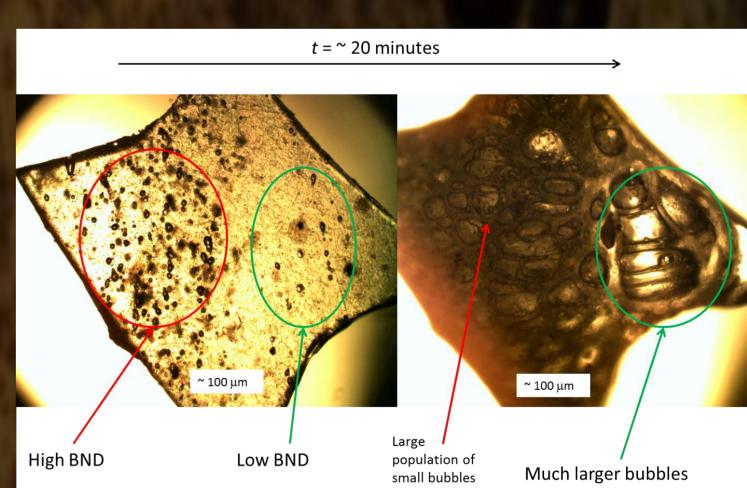


Fig 6. Contrasting initial and final bubble size distributions, the hold temperature here is 775°C

4. Further findings and conclusions

- Water content strongly influences vesicle growth rates, which are ~7 times higher in the water-rich Chaitén rhyolite than the GOVC peralkaline rhyolite (0.14 wt%) used in [1].
- We estimate bubble nucleation rates (J) of >1.5 x 10^{10} m⁻³ s⁻¹, from change in bubble number through time. This matches the lower end of J values from decompression experiments[6]. High nucleation rates occurred for ~30 seconds, prior to bubble number reduction due to coalescence during foaming.
- Bubble growth rates in our experiments were poorly represented by the Navon et al ('98) model, with worst fits at low temperature (high viscosity). However, we found a good correlation between measured growth rates and modelled diffusivity, implying that volatile diffusivity may have been a limiting factor.

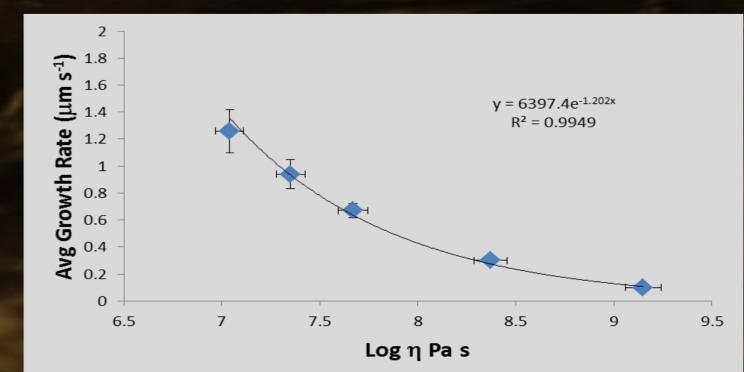


Fig 7. Bubble growth rate (V_r) increases with decreasing melt viscosity(η) where: $V_r \approx \exp(-1.202 \, \eta)$.

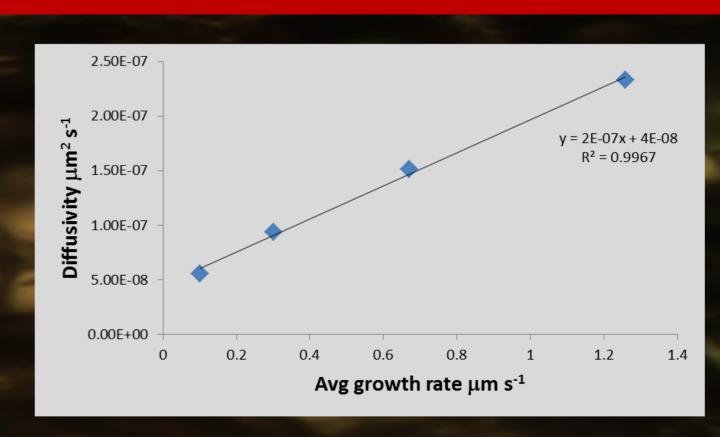


Fig 8. At higher temperatures (> 725°C) there is a positive linear relationship between bubble growth rate and modelled diffusivity.

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⁴Zhang, Y (1999) H₂O in rhyolitic glasses and melts: Measurements, speciation, solubility, and diffusion. Reviews in Geophysics. ⁵Navon et al. (1998) Bubble growth in highly viscous melts: theory, experiments, and autoexplosivity of dome lavas. *EPSL*. 160, 763-776. ⁶Cluzel N. et al. (2008) Contrib Mineral Petrol (2008) 156:745–763; Gardner & Ketcham (2011) Contrib Mineral Petrol 162:929–943. ⁷Giordano, D; Russell, J.K; Dingwell, D.B. (2008) Viscosity of magmatic liquids: A model. Earth and planetary science letters. 271, 123-134